

The Price-Weller-Pinkel Dynamical Instability Model

The Price-Weller-Pinkel (PWP) vertical mixing model (Price *et al.*, 1986) has been adapted to provide vertical mixing in HYCOM. The PWP algorithm is executed at p grid points where thermodynamical variables are mixed. To perform these calculations, momentum components are horizontally interpolated to the p points. After the thermodynamical variables have been mixed at the p points, a special algorithm described below is executed to perform momentum mixing at the u, v grid points.

The first step is to apply the surface boundary conditions. Surface momentum fluxes are applied when the momentum equation is solved, with the surface fluxes acting to accelerate the water in model layer 1 only. When the PWP subroutine is called, the surface fluxes of heat and mass are distributed over model layer 1 except for penetrating shortwave radiation, which is distributed over multiple layers. If static instability is present near the ocean surface after applying the surface fluxes, it is relieved as follows: If model layer 2 is less dense than layer 1, the water in both layers is completely mixed. If the water in layer 3 is less dense than this mixed water, it is completely mixed with the water from layers 1 and 2. This process is repeated until a model layer is encountered that is denser than the mixed water above.

An initial diagnosis of mixed layer depth is performed next. Moving down from the surface, the mixed layer depth is set to the depth of the first interface across which the density jump exceeds a prescribed value. The mixed layer base always resides on a model interface in the PWP model. All model variables on p grid points, including the interpolated momentum components, are homogenized within the mixed layer.

To perform bulk Richardson number mixing, this number is calculated using

$$R_b = \frac{g \Delta \rho h}{\rho_0 [(\Delta u)^2 + (\Delta v)^2]}, \quad (1)$$

where h is the diagnosed mixed layer thickness. Density and velocity differences are taken between the mixed layer and the model layer immediately below. If $R_b < 0.65$, the mixed layer entrains the layer immediately below it and all variables are homogenized within the new mixed layer. The process is then repeated, with the mixed layer entraining additional layers, until $R_b \geq 0.65$.

To perform gradient Richardson number mixing, this number is estimated using

$$R_g = \frac{g \Delta \rho \Delta z}{\rho_0 [(\Delta u)^2 + (\Delta v)^2]}, \quad (2)$$

and mixing is performed if $R_g < 0.25$. At each interface, $\Delta \rho$, Δu , and Δv are estimated as the difference between the layer above and the layer below, while Δz is estimated as the average thickness of the layers above and below.

The following procedure is used to perform the shear instability mixing. The number R_g is estimated at all vertical interfaces between the one upon which the mixed layer base resides and the bottom. The interface with the smallest value of R_g is identified. If this value is less than 0.25, the layers above and below the interface are partially mixed so that the value of R_g is increased to 0.30. New values of R_g are calculated, and then the new interface with the smallest

value of R_g is identified. If this value is less than 0.25, the two adjacent layers are partially mixed in the same manner. This process is repeated until the minimum value of R_g over all layers exceeds 0.25.

After this mixing process has been completed, the depth of the mixed layer base is again diagnosed as the depth of the first interface below the surface where the density jump exceeds a prescribed value. The final vertical homogenization is performed only for thermodynamical variables and other scalars that are stored at p grid points. After completing the mixing process at pressure grid points, the depth of the mixed layer base is horizontally interpolated to u and v grid points, and the velocity components are homogenized within the mixed layer. On both sets of grids, the model interface closest to the interpolated mixed layer depth is identified as the mixed layer base, and vertical homogenization is performed between the surface and the identified interface. Although the PWP algorithm provides for shear instability mixing beneath the surface mixed layer, it does not provide for background mixing due to other processes such as internal wave breaking. When the PWP model is selected, the hybrid explicit (MICOM-like) diapycnal-mixing algorithm is also activated to provide this additional mixing.

REFERENCE

Price, J. F., R. A. Weller, and R. Pinkel, 1986: Diurnal cycling: Observations and models of the upper ocean response to diurnal heating, cooling, and wind mixing. *J. Geophys. Res.*, **91**, 8411-8427.